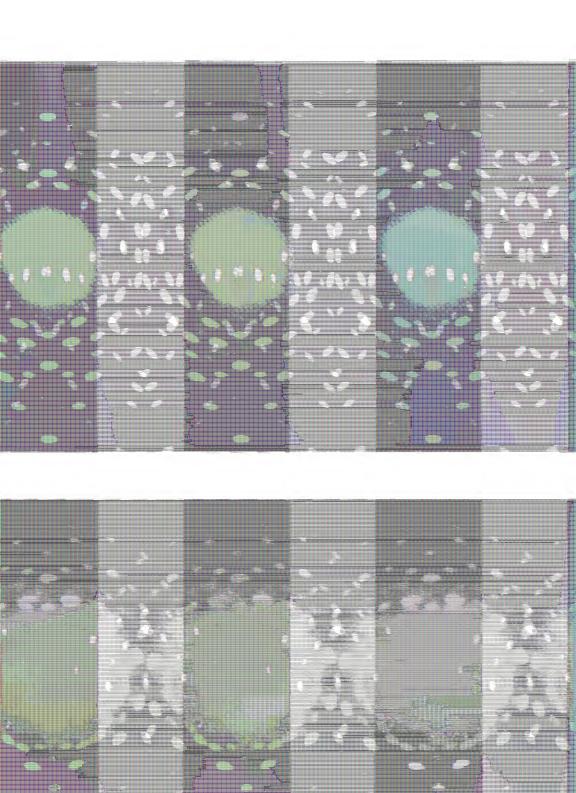
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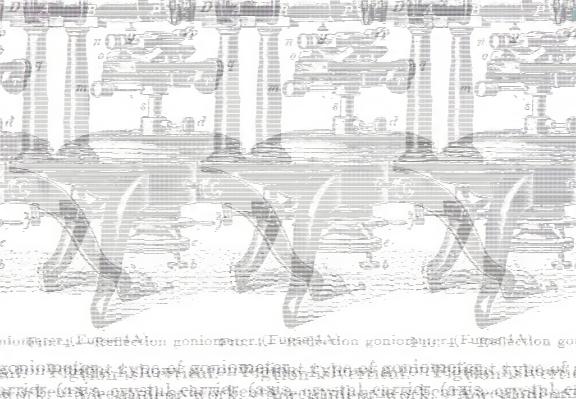
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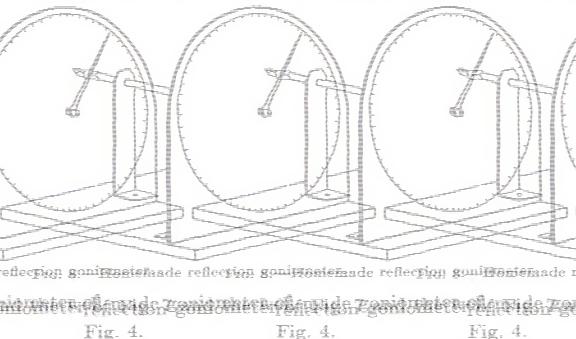
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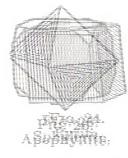
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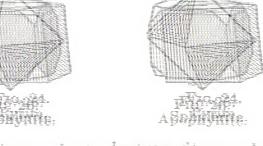
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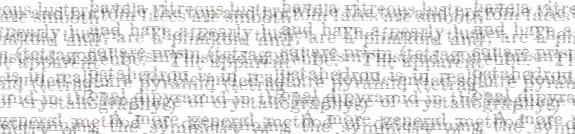
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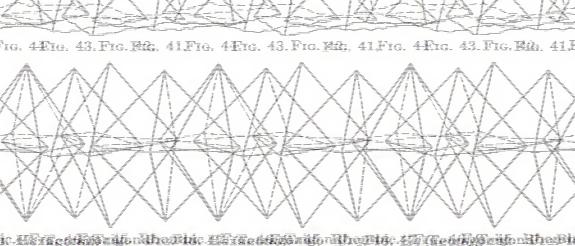
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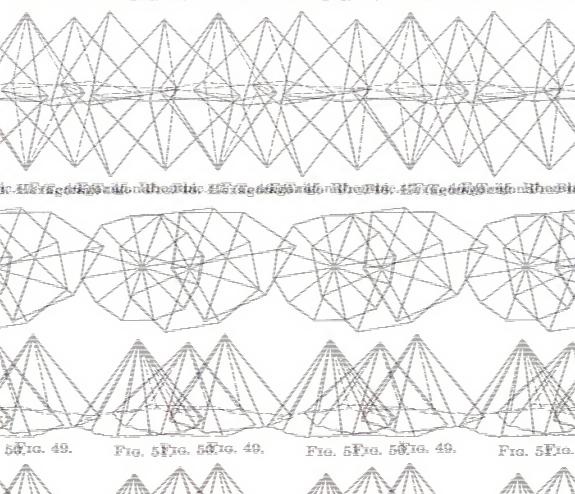
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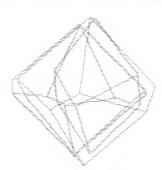




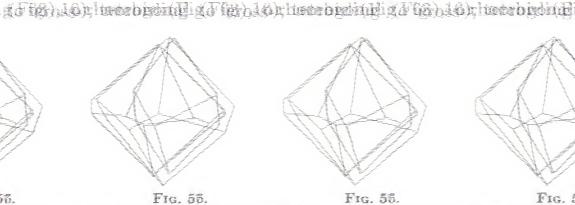
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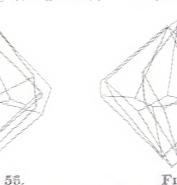
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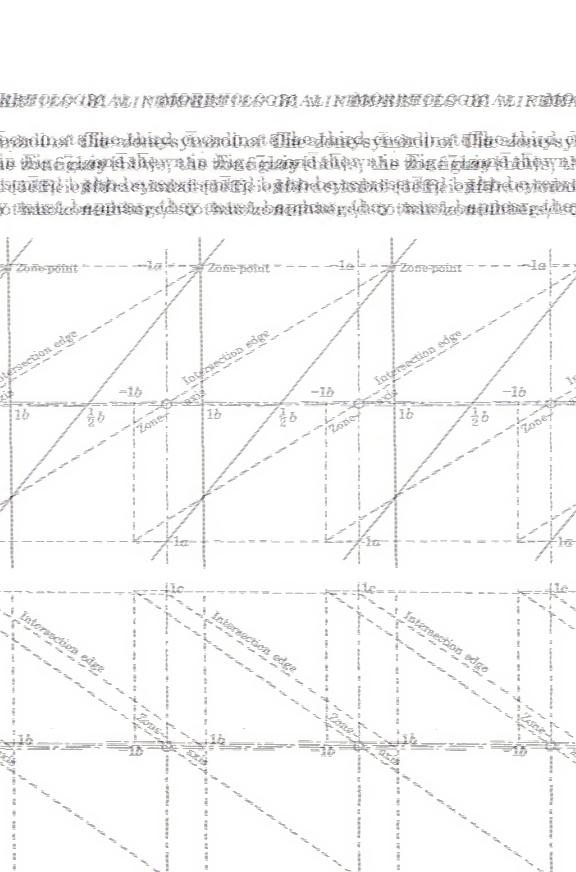
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projection of any face is the intersection of a line drawn from the south pole of the sphere (S in the drawing) to the pole of the face with the plane of projection. The face e, for example, is projected at e'. Faces on the upper half of the crystal are projected within the primitive circle, those on the lower half of the crystal are projected outside the primitive circle, and faces of the vertical zone, such as a, m, b, appear on the primitive circle.

Figure 89a, which represents a vertical section of a crystal within a sphere, shows why faces on the lower half of the crystal

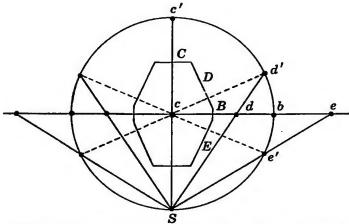


Fig. 89a.—Method of making a stereographic projection of the faces of a vertical zone.

are projected outside the primitive circle. The pole of the face E, for example, is at e' and its projection at e.

The important property of the stereographic projection is that all circles on the sphere appear as circles or as straight lines (circles of infinite radius) on the projection. Circles on the sphere are of two kinds: great circles, whose planes pass through the center of the sphere and small circles, whose planes do not pass through the center of the sphere. All vertical great circles appear as diameters of the primitive circle. All oblique great circles are arcs of great circles passing through extremities of diameters.

The method of projecting poles on the projection depends upon their position. Poles on the equatorial circle are laid off directly on the primitive. Poles on vertical great circles are projected by first laying off the angles on the primitive. Lines tra hardprovide toperint a hardprovide polarita hardppositie polein generalise poleinta hardppositie polein generalise poleinta hardppositie polein generalise politica politi

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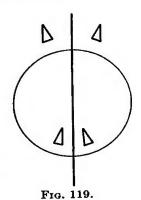
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List	of	Forms	in	the	Prismatic	Class
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Rhombic prisms Rhombic prism Rhombic prism Pinakoids Pinakoid Pinakoid Pinakoid	4 faces 4 faces 4 faces 2 faces 2 faces 2 faces 2 faces 2 faces	{hkl}, {0kl} {hk0} {h0l}, {100} {010} {001}	
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Rhombic Prisms {hkl}, {hkl}. These two forms occur independently, but together they constitute a figure that resembles a pyramid; hence the name *hemipyramid* is sometimes used.

ed. Figure 120

represents an {hkl} form.

Rhombic Prism {Okl}. An open form consisting of four faces, each parallel to the a-axis. The a-axis is sometimes called the clino-axis, Fig. 121.

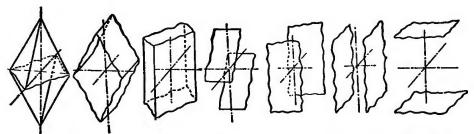


Fig. 120. Fig. 121. Fig. 122. Fig. 123. Fig. 124. Fig. 125. Fig. 126. Figs. 120-126.—The seven type forms in the prismatic class.

Rhombic Prism {hk0}. An open form consisting of four faces each parallel to the vertical axis, Fig. 122.

Pinakoids $\{h0l\}$, $\{\overline{h}0l\}$. These forms, each composed of two opposite parallel faces parallel to the *b*-axis (often called the ortho-axis), are independent of each other. Figure 123 represents $\{h0l\}$.

Pinakoid {100}. This form may be called the front pinakoid, but it is also known as the orthopinakoid, Fig. 124.

Pinakoid {010}. This may be called the side pinakoid, but it is also known as the clinopinakoid, Fig. 125.

Pinakoid {001}. This form is usually known as the basal pinakoid, but its faces are inclined and not perpendicular to the c-axis, Fig. 126.

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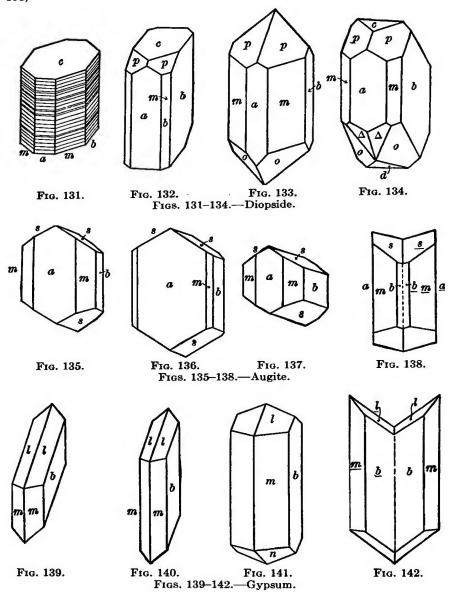
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of augite crystals. Twins with {100} as twin-plane are common (see Fig. 138).



Gypsum. $\hat{a}: \bar{b}: \hat{c} = 0.690:1:0.412$; $\angle \beta = 80^{\circ} 42'$. Usual forms: $m\{110\}$, $l\{111\}$, $b\{010\}$, $n\{\bar{1}11\}$, $e\{\bar{1}03\}$. Cleavage parallel to b. Interfacial angles: $mm(110:1\bar{1}0) = 68^{\circ} 30'$; $bm(010:110) = 55^{\circ} 45'$; $ll(111:1\bar{1}1) = 36^{\circ} 12'$; $bn(010:\bar{1}11) = 69^{\circ} 20'$; $bl(010:111) = 71^{\circ} 54'$; $be(010:\bar{1}03) = 90^{\circ} 0'$;

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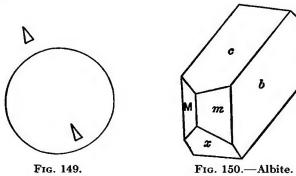
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Figure 149 is the stereographic projection of the general form of this class.

List of Forms in the Pinakoidal Class {hkl}, {hkl}, {hkl}, {hkl} Pinakoids 2 faces 2 faces $\{h01\}, \{\bar{h}01\}$ Pinakoids Pinakoids 2 faces $\{0kl\}, \{0\bar{k}l\}$ Pinakoids 2 faces |hk0|, |hk0| [100] Pinakoid 2 faces {010} Pinakoid 2 faces 2 faces {001} Pinakoid

All forms are pinakoids, each of which consists of two opposite parallel faces.



Combinations. The appearance of triclinic crystals depends largely upon the obliquity of the axes. Many of them closely approach monoclinic crystals in angles. This is especially the case with the plagioclase feldspars.

Examples

Comparatively few minerals crystallize in this class. The only common ones are the plagioclase feldspars, rhodonite, kyanite, and microcline. Albite (NaAlSi₂O₃) is selected as the best mineral for study. Albite crystals are usually so small that measurements must be made by the reflection goniometer. Microcline is triclinic but is so close to orthoclase in angles that it may pass for monoclinic. (The optical properties, especially the oblique extinction on the 001 cleavage face, prove that it is triclinic.)

Albite. $\check{a}:\bar{b}:\dot{c}=0.633:1:0.556;\ \angle\alpha=93^\circ\ 58';\ \angle\beta=63^\circ\ 39';\ \angle\gamma=87^\circ\ 31'.$ Usual forms; $m\{110\},\ M\{1\overline{1}0\},\ c\{001\},\ b\{010\},\ x\{\overline{1}01\},\ y\{\overline{2}01\},\ f\{130\},\ z\{1\overline{3}0\},\ n\{021\},\ p\{\overline{1}11\},\ o\{\overline{1}\overline{1}1\}.$ Cleavage parallel to c and b. Interfacial angles: $mM(110:1\overline{1}0)=59^\circ\ 16\frac{1}{2}';\ mf(110:130)=30^\circ\ 24';\ Mz(1\overline{1}0:1\overline{3}0)=29^\circ\ 36';\ mb(110:010)=60^\circ\ 58';\ cx(001:\overline{1}01)=52^\circ\ 6\frac{1}{2}';\ cy(001:\overline{2}01)=81^\circ\ 53';$ Figure 150 represents an albite crystal with the forms; $c\{001\},\ b\{010\},\ m\{110\},\ M\{1\overline{1}0\}$ and $x\{\overline{1}01\}.$

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the symbols of f, z, y, and p. (c = 001; b = 010). The determinations are made just as they were in the case of Fig. 147; the only difference is that the b-axis is not normal to the b(010) face. The following symbols are obtained f = (130); $z = (\overline{130}); y = (\overline{201}); p = (\overline{111}).$

16. TETRAGONAL SYSTEM

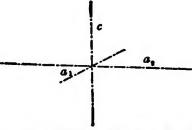
The tetragonal system includes all crystals with a single tetragonal zone.

In the tetragonal system there are seven crystal classes with varying symmetry, but they all have either a single 4-fold axis of symmetry (A₄) or a single 4-fold rotatory-reflection axis of symmetry (P4).

There are three axes of reference at right angles to each other, two of which are interchangeable. The axes are designated a₁:a₂:c. The unit lengths on a₁ and a₂ each equal to unity and the unit length on c in general either greater or less than unity. The axes of reference for the tetragonal system are represented in Fig. 154. The A4 or P4 axis of symmetry is taken as the c-axis and the lateral axes of reference are either 2-fold axes of

symmetry, lines in planes of symmetry, or lines parallel to prominent edges.

Since the lateral axes a_1 and a_2 are interchangeable, interfacial angles in the prism zone are the same for corresponding faces of all tetragonal crystals. The angles for some of the more com- Fig. 154.—Tetragonal axes of refmon forms are: $(100:110) = 45^{\circ}$



erence for zircon.

0'; $(100:320) = 33^{\circ} 41'$; $(100:210) = 26^{\circ} 34'$; $(100:310) = 18^{\circ}$ 26'; (100:410) = 14° 2'.

(15) Ditetragonal Dipyramidal Class. A4.4A2.5P.C

(Holosymmetric)

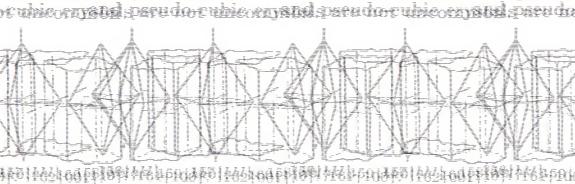
The 2-fold axes are normal to the 4-fold axis. Four vertical planes of symmetry intersect each other at angles of 45°, and the fifth is normal to these four.

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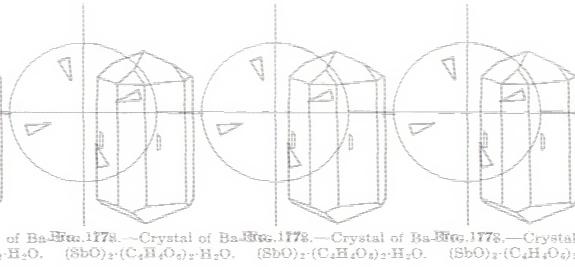
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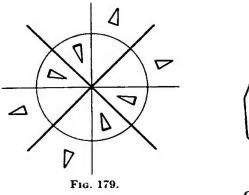
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Tetragonal scalenohedrons	8 faces	{hkl}, {hkl}
Tetragonal disphenoids	4 faces	{hhl}, {hhl}
Tetragonal dipyramid	8 faces	{ h0 1}
Ditetragonal prism	8 faces	{hk0}
Tetragonal prisms	4 faces	{110}, {100}
Pinakoid	2 faces	{001}

The two correlative scalenohedrons are congruent forms distinguished as right $\{hkl\}$ and left $\{hkl\}$. The right-handed form is shown in stereographic projection in Fig. 179.



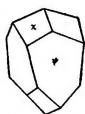


Fig. 180.— Chalcopyrite crystal.

Chalcopyrite is the best known representative of this class. Figure 180 illustrates a crystal with the tetragonal disphenoid $\psi(772)$ and the tetragonal scalenohedron $\chi(122)$ from French Creek, Pa., described by Penfield.

(12) Tetragonal Trapezohedral Class. A.4A.

(Holoaxial hemisymmetric)

The 4-fold axis of symmetry is taken as the c-axis and either pair of 2-fold axes at right angles to each other as the lateral axes a_1 and a_2 . The type forms are:

Tetragonal trapezohedrons	8 faces	[hkl], {khl}
Tetragonal dipyramids	8 faces	h01}, {hh1}
Ditetragonal prism	8 faces	{hk0}
Tetragonal prisms	4 faces	110 , [100]
Pinakoid	2 faces	1001

The two correlative trapezohedrons are enantiomorphs. They are distinguished as right [hkl] and left [khl]. Figure 181 is a stereographic projection of the general form.

The best example of this class is methyl ammonium iodid, NH₁(CH₁)I, described by Wagner. Figure 182 illustrates a crystal with the tetragonal prism $a|100\rangle$, the pinakoid $c|001\rangle$, and the tetragonal trapezohedron $t|166\rangle$.

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form, modified by the unit dipyramid $p\{111\}$ and the left dipyramid $s\{131\}$, which is a general form.

(14) Ditetragonal Pyramidal Class. A₄-4P

(Polar hemisymmetric)

The 4-fold axis of symmetry is the intersection of the four planes of symmetry. The 4-fold axis is the c-axis; the lateral axes a_1 and a_2 are in either pair of symmetry planes (which are mutually perpendicular) and at right angles to the c-axis. The type forms are:

 Ditetragonal pyramids
 8 faces
 {hkl}, {hkl}

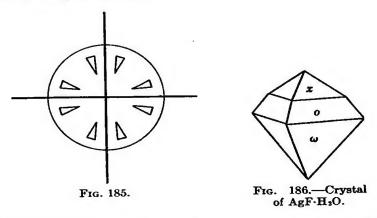
 Tetragonal pyramids
 4 faces
 {hhl}, {hhl}, {h0l}, {h0l}

 Ditetragonal prism
 8 faces
 {hk0}

 Tetragonal prisms
 4 faces
 {110}, {100}

 Pedions
 1 face
 {001}, {001}

The correlative ditetragonal pyramids are congruent and may be distinguished as upper $\{hkl\}$, and lower $\{hkl\}$. Figure 185 is a stereographic projection of the general form.



One of the best examples of this class is hydrous silver fluorid, AgF·H₂O, represented by Fig. 186. Here the forms are $x\{113\}$, $o\{111\}$, and $\omega\{11\overline{1}\}$, all tetragonal pyramids.

Graphic Determination of Indices and Axial Ratio in the Tetragonal System. A plan and elevation of a zircon crystal are shown in Fig. 187. The unit dipyramid $\{111\}$ is the p-face; the problem is to determine the symbol of u and the axial ratio a:c. In the elevation, lines parallel to the projections of p and u are drawn through the point x to intersect the c-axis. Since the distance os is equal to 3 times the distance or, the symbol of u is $1a_1:1a_2:3c$ or (331). The distance or is equal to about 0.64 of the

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(27) Dihexagonal Dipyramidal Class. A₆·6A₂·7P·C

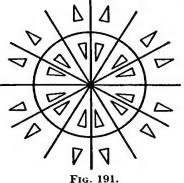
(Holosymmetric)

The 2-fold axes are normal to the 6-fold axis. There are six vertical planes of symmetry at angles of 30° apart. The other plane of symmetry is perpendicular to these six.

List of Forms in the Dihexagonal Dipyramidal Class

Dihexagonal dipyramid	24 faces	{hkil}
Hexagonal dipyramid	12 faces	$\{\mathbf{h}\cdot\mathbf{h}\cdot\overline{\mathbf{2h}}\cdot\mathbf{l}\}$
Hexagonal dipyramid	12 faces	{ h0 h1}
Dihexagonal prism	12 faces	{hki0}
Hexagonal prism	6 faces	$\{11\overline{2}0\}$
Hexagonal prism	6 faces	{1010}
Pinakoid	2 faces	{0001}
(In the above sur	mbols h > 1	k)

The stereographic projection of the general form of this class is shown in Fig. 191.



This form consists of 24 faces (scalene triangles in the ideal form), each of which cuts the four axes at unequal distances. The symbols of the upper faces of the dihexagonal

Dihexagonal Dipyramid {hkil}.

dipyramid are given in Fig. 192. The angles over alternate polar edges are equal. (Fig. 193.)

Hexagonal Dipyramid $\{h \cdot h \cdot \overline{2h} \cdot l\}$ (Pyramid of the second order).

The faces cut two of the lateral axes at equal but greater distances than the third lateral axis. (Fig. 194.) This form differs from {h0h1} only in position.

Hexagonal Dipyramid {h0h1} (Pyramid of the first order). The faces cut two of the lateral axes, but are parallel to the third. (Fig. 195.)

Dihexagonal Prism {hki0}. All the faces are in a vertical zone, each being parallel to the vertical axis. Alternate angles are equal. (Fig. 196.)

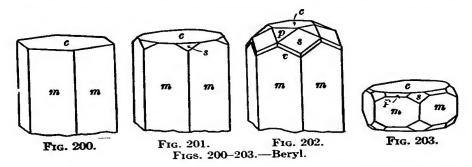
Hexagonal Prism {1120} (Prism of the second order). This form is similar to {1010} except in position. (Fig. 197.)

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Combinations. The habit is prismatic, pyramidal, or tabular. Simple combinations are the rule in this class. As beryl is the only common mineral belonging to this class, it is the only example given for practice.

Example

Beryl. $\dot{c}=0.498$. Usual forms: $c\{0001\}$, $m\{10\overline{1}0\}$, $p\{10\overline{1}1\}$, $s\{11\overline{2}1\}$, $v\{21\overline{3}1\}$. Interfacial angles: $mm(10\overline{1}0:01\overline{1}0)=60^{\circ}$ 0'; $cs(0001:11\overline{2}1)=44^{\circ}$ 56'; $cp(0001:10\overline{1}1)=29^{\circ}$ 57'; $mv(10\overline{1}0:21\overline{3}1)=37^{\circ}$ 49'; $ms(10\overline{1}0:11\overline{2}1)=52^{\circ}$ 17'. Figures 200 and 201 are the ordinary combinations. Figure 202



has in addition the general form $v\{21\overline{3}1\}$. Figure 203 represents beryl of tabular habit, which is rare as compared with the prismatic habit.

(21) Trigonal Dipyramidal Class. CA₆(A₃)(P)

(Trigonal tetartosymmetric)

The plane of symmetry is normal to the axis of symmetry which is taken as the c-axis. The lateral axes of reference should be taken parallel to prominent edges or faces, but in the plane of symmetry. The type forms are:

```
\{hkil\}, \{ikhl\}, \{khil\}, \{kihl\}
                              6 faces
Trigonal dipyramids
                                           \{\mathbf{h}\cdot\mathbf{h}\cdot\mathbf{2h}\cdot\mathbf{1}\},\{\mathbf{2h}\cdot\mathbf{h}\cdot\mathbf{h}\cdot\mathbf{1}\}
Trigonal dipyramids
                              6 faces
                                           {h0h1}, {0hh1}
Trigonal dipyramids
                              6 faces
                                           |hki0|, {ikh0}, {hki0}, {kih0}
                              3 faces
Trigonal prisms
                                           {1120}, {2110}, {1010}, {0110}
                              3 faces
Trigonal prisms
                                           10001
                              2 faces
Pinakoid
                        (In the above symbols h > k)
```

There are four correlative dipyramids possible for any given values of h, k, and l. Figure 204 is a stereographic projection of the general form.

No representative of this class is known. The forms given above may be deduced from the symmetry, but no drawing of a crystal of this class with the general form can be made, since no axial ratio a:c is known.

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The only certainly known example of this class is benitoite, BaTiSi2O2. from San Benito County, Calif., described by Louderback in 1907. Figure 206 represents a typical crystal with the forms: trigonal prism $m\{10\overline{10}\}$, trigonal prism $\mu\{01\overline{1}0\}$, trigonal dipyramid $p\{10\overline{1}1\}$, trigonal dipyramid $\pi\{01\overline{1}1\}$, and pinakoid $c\{0001\}$.

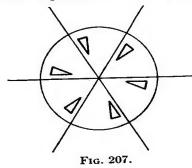
(23) Hexagonal Pyramidal Class. A6

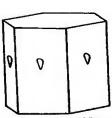
(Polar hexagonal tetartosymmetric)

The axis of symmetry is the c-axis; the lateral axes of reference are parallel to prominent edges and also normal to the c-axis. The type forms are:

```
{hkīl}, {ikhl}, {hkīl}, {ikhl}
                                                    6 faces
Hexagonal pyramids
                                                                          \{\mathbf{h}\cdot\mathbf{h}\cdot\mathbf{2}\overline{\mathbf{h}}\cdot\mathbf{1}\}, \{\mathbf{h}\cdot\mathbf{h}\cdot\overline{\mathbf{2}}\overline{\mathbf{h}}\cdot\overline{\mathbf{1}}\}, \{\mathbf{h}\mathbf{0}\overline{\mathbf{h}}\mathbf{1}\}, \{\mathbf{h}\mathbf{0}\overline{\mathbf{h}}\overline{\mathbf{1}}\}
Hexagonal pyramids
                                                    6 faces
                                                                          \{hki0\}, \{i\overline{kh}0\}, \{11\overline{2}0, \{10\overline{1}0\}\}
Hexagonal prisms
                                                    6 faces
                                                     2 faces
                                                                          \{0001\}, \{0001\}
Pedions
                                               (In the above symbols h > k)
```

The four correlative general forms of the first line are distinguished in order as upper right, upper left, lower right, and lower left pyramids in the Figure 207 is a stereographic projection of the general form. order given.





F 1 G . 208.-Nepheline crystal with etch-figures.

Nepheline (essentially NaAlSiO4) has been assigned to this class on the basis of etch-figures (see Fig. 208), where the forms are the unit hexagonal prism $\{10\overline{10}\}$, and the upper and lower pedions $\{0001\}$ and $\{000\overline{1}\}$.

(24) Hexagonal Trapezohedral Class. As-6A2

(Holoaxial hemisymmetric)

The 2-fold axes of symmetry are normal to the 6-fold axis, which is taken as the c-axis. The lateral axes of reference are either of two sets of the 2-fold axes at 120° to each other. The type forms are:

Hexagonal trapezohedrons Hexagonal dipyramids Dihexagonal prisms Hexagonal prisms Pinakoid	2 faces	•
(T- the abone	OMETHINGS /L	- N 1

(In the above symbols h > k)

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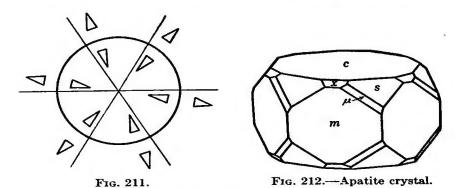
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O.—Crystap. 389.Fig. 210.—Crystap. 289.Fig. 210.—Crystap. 289.Fig. 210.—Crystap. 389.Fig. 389

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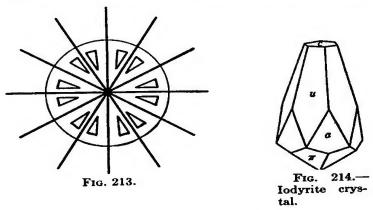
(26) Dihexagonal Pyramidal Class. A6.6P

(Polar hemisymmetric)

The six vertical planes of symmetry intersect in the six-fold axis of symmetry which is chosen for the c-axis. The lateral axes are one of two sets of lines normal to the c-axis, but within the planes of symmetry. The type forms are:

12 faces {hkīl}, {hkil} Dihexagonal pyramids $\{h \cdot h \cdot 2\overline{h} \cdot 1\}, \{h0\overline{h}1\}$ Hexagonal pyramids 6 faces $\{\mathbf{h}\cdot\mathbf{h}\cdot\overline{\mathbf{2h}}\cdot\overline{\mathbf{l}}\},\ \{\mathbf{h}0\overline{\mathbf{h}}\overline{\mathbf{l}}\}$ Hexagonal pyramids 6 faces |hki0| Dihexagonal prism 12 faces $\{10\overline{1}0\}, \{11\overline{2}0\}$ 6 faces Hexagonal prisms {0001}, {0001} 1 face **Pedions** (In the above symbols h > k)

The two correlative general forms given on the first line are distinguished as upper and lower. The stereographic projection is shown in Fig. 213.



Iodyrite, AgI, furnishes a good example of this class (see Fig. 214). Here the forms are the upper pedion $c\{0001\}$, upper hexagonal pyramid $u\{40\overline{4}1\}$, hexagonal prism $a\{11\overline{2}0\}$, and lower hexagonal pyramid $\pi\{40\overline{4}5\}$.

Ice also probably belongs to this class. (See p. 326.)

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where f=2p+2q-r, where f=2p+2q-r, where f=2p+2q-r, where f=2p-q+2r, r=2p-q+2r, where r=2p+q. (where r=2p+q.) Directly r=2p-q+2r, r

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Calcite. c = 0.854 ($\alpha = 101^{\circ}$ 54' for the three-axis method). Usual forms: $c\{0001\} = \{111\}$, $m\{10\overline{1}0\} = \{2\overline{1}1\}$, $a\{11\overline{2}0\} = \{10\overline{1}\}$, $e\{01\overline{1}2\} = \{110\}$, $r\{10\overline{1}1\} = \{100\}$, $f\{02\overline{2}1\} = \{11\overline{1}\}$, $h\{03\overline{3}2\} = \{55\overline{4}\}$, $M\{40\overline{4}1\} = \{3\overline{1}1\}$, $v\{21\overline{3}1\} = \{20\overline{1}\}$, $y\{32\overline{5}1\} = \{30\overline{2}\}$, $t\{21\overline{3}4\} = \{310\}$. Cleavage, perfect parallel to r. Interfacial angles: $ee(01\overline{1}2:\overline{1}012) = 45^{\circ}$ 3'; $em(01\overline{1}2:01\overline{1}0) = 63^{\circ}$ 45'; $rr(10\overline{1}1:\overline{1}101) = 74^{\circ}$ 55'; $rm(10\overline{1}1:10\overline{1}0) = 45^{\circ}$ 23½'; $ff(02\overline{2}1:\overline{2}021) = 101^{\circ}$ 9'; $fm(02\overline{2}1:01\overline{1}0) = 26^{\circ}$ 53'; $MM(40\overline{4}1:\overline{4}401) = 114^{\circ}$ 10'; $Mm(40\overline{4}1:10\overline{1}0) = 14^{\circ}$ 13'; $hh(03\overline{3}2:\overline{3}\overline{3}02) = 91^{\circ}$ 42'; $vv(21\overline{3}1:\overline{2}3\overline{1}1) = 75^{\circ}$ 22'; $vv(21\overline{3}1:3\overline{1}21) = 35^{\circ}$ 36'; $vv(21\overline{3}1:12\overline{3}1) = 47^{\circ}$ 1'; $yy(32\overline{5}1:\overline{3}5\overline{2}1) = 70^{\circ}$ 59'; $yy(32\overline{5}1:5\overline{2}31) = 45^{\circ}$ 32'; $vy(21\overline{3}1:32\overline{5}1) = 8^{\circ}$ 53'; $vv(10\overline{1}1:21\overline{3}1) = 29^{\circ}$ 1½'; $mv(10\overline{1}0:21\overline{3}1) = 28^{\circ}$ 4'; $tv(21\overline{3}4:3\overline{1}24) = 20^{\circ}$ 36½'; $tv(21\overline{3}4:01\overline{1}2) = 20^{\circ}$ 57½'.

Figures 220 to 231 represent some of the common types of calcite crystals. The dotted lines in the figures represent cleavage planes which aid in distinguishing positive and negative forms.

(16) Trigonal Pyramidal Class. A.

(Ogdosymmetric)

The c-axis is the 3-fold axis of symmetry; the lateral axes of reference a_1 , a_2 , and a_3 are taken parallel to prominent edges. The type forms are:

```
{pqr}, {prq}.... Trigonal pyramids 3 faces
                                                                       |hkil|, |ikhl|
{efg}, {feg}.... Trigonal pyramids 3 faces (where e = 2p - q + 2r,
                                                                       |kihl|, |khil|
    f=2p+2q-r,
  g = -p + 2q + 2r
        {gēf}, {gfe}.... Trigonal pyramids
                                                           3 faces
                                                                       hkil, ikhl
        {qrp}, {rqp}.... Trigonal pyramids
                                                           3 faces
                                                                       {kihl}, {khil}
        {pqr}, {prq}.... Trigonal pyramids
                                                           3 faces
                                                                       \{\mathbf{h}\cdot\mathbf{h}\cdot\overline{2}\mathbf{h}\cdot\mathbf{l}\}, \{2\mathbf{h}\cdot\overline{\mathbf{h}}\cdot\overline{\mathbf{h}}\cdot\mathbf{l}\}
        {rqp}, {qrp}....
                               Trigonal pyramids
                                                           3 faces
                                                                       \{\mathbf{h}\cdot\mathbf{h}\cdot\mathbf{2h}\cdot\mathbf{1}\}, \{\mathbf{2h}\cdot\mathbf{h}\cdot\mathbf{h}\cdot\mathbf{1}\}
   (where p + r = 2q)
                                                                       {h0h1}, {0hh1}
        {pqq}, {eef} ... Trigonal pyramids
                                                           3 faces
        {fēē}, {qqp}.... Trigonal pyramids
                                                           3 faces
                                                                       hohi . ohhi
   (where e=2p+q,
      f = -p + 4q)
        {pqr}, {prq}.... Trigonal prisms
                                                           3 faces
                                                                       |hki0|, |ihk0|
        {qrp}, {rqp}.... Trigonal prisms
                                                                       {kih0}, |khi0}
                                                           3 faces
       {101}, {110}.... Trigonal prisms
                                                                       1120|, |2110|
                                                           3 faces
       \{2\overline{11}\}, \{11\overline{2}\}\dots Trigonal prisms
                                                           3 faces
                                                                       \{10\overline{1}0\}, \{0\overline{1}\overline{1}0\}
       {111}, {111}.... Pedions
                                                           1 face
                                                                       \{0001\}, \{000\overline{1}\}
                              (In the above symbols h > k)
```

There are no less than eight correlative trigonal pyramids possible as general forms for any given values of h, k, and l. These are named in the order listed as upper right positive, upper left positive, upper right negative, upper left negative, and then the same order for the lower forms. The other correlative forms are named in an analogous manner with the appropriate omissions.

Figure 232 is a stereographic projection of the general form.

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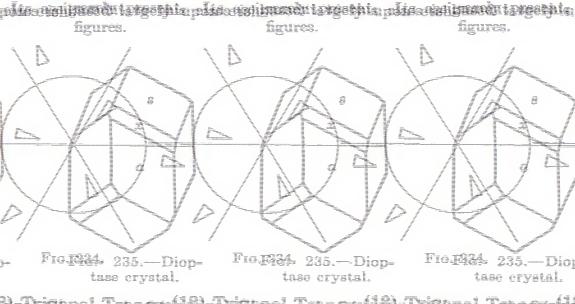
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 $J = 2p + 2q + \tilde{r}.$

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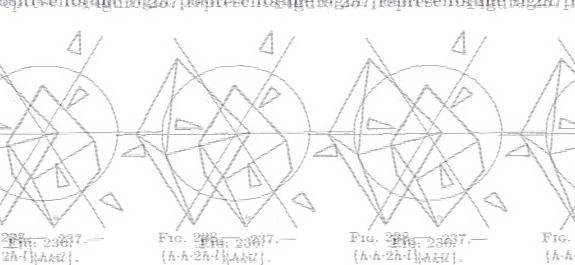
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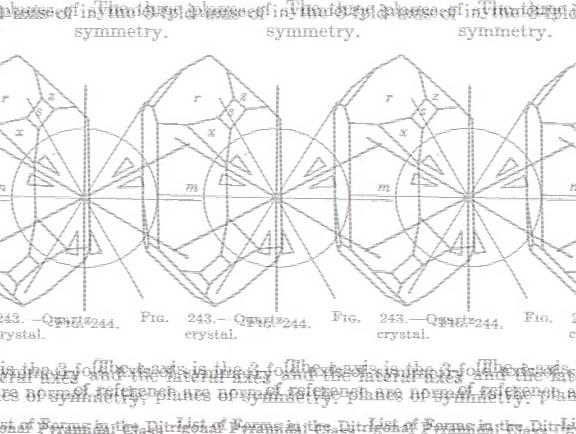
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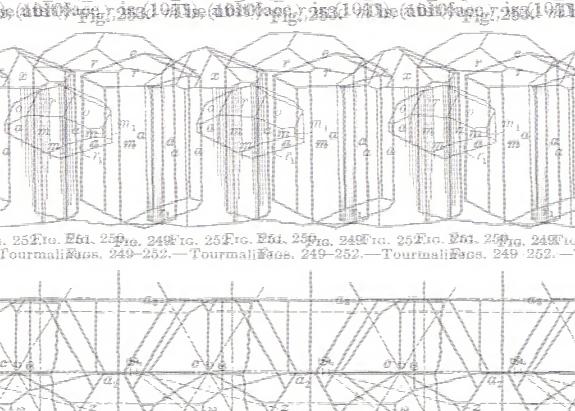
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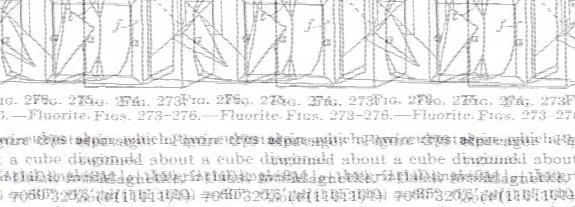
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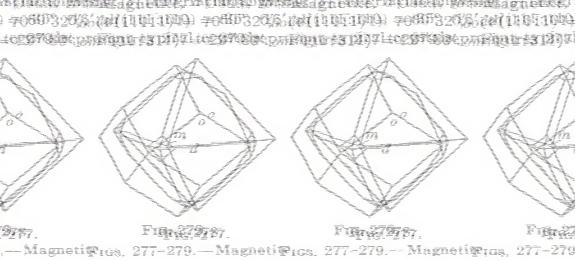
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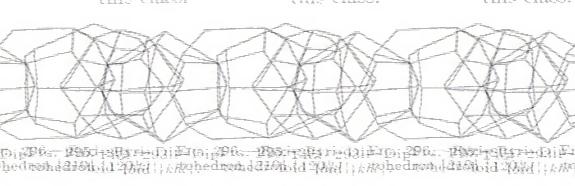
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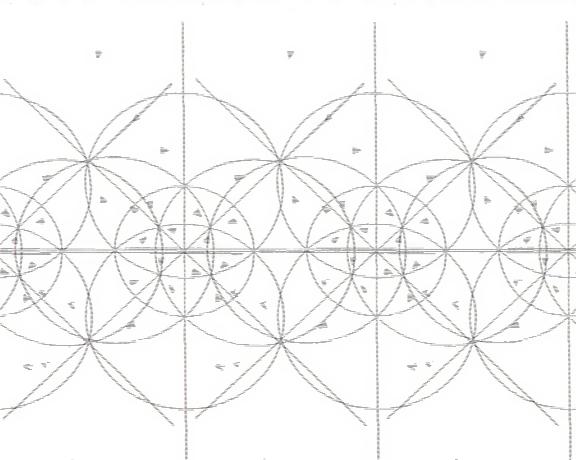


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Barium nitrate is a good example of the class. Figure 310 represents a crystal with the forms: the cube $a\{100\}$, tetrahedrons $a\{111\}$ and $a_1\{1\overline{1}1\}$. and the tetartoid, $\lambda\{421\}$.

 α_3

pyrite crystal.

It is not necessary to find a tetartoid on a crystal in order to assign it. to this class. The presence of both a tetrahedron and a pyritohedron on a crystal practically proves that it belongs to this class.

Graphic Determination Indices in the Isometric System. There are no axial ratios to be determined but simply the indices of the faces. A plan and elevation are made from the interfacial angles. Figure 311 represents pyrite with the faces a, e, and o. The intercepts of the e-face in terms of the unit Fig. 311.—Plan and elevation of a are seen to be 1a1:2a2: ∞a2 which reduces to the Miller

Fig. 312.—Octa-

symbol (210). The intercepts of e_1 are: $\alpha_1:\frac{1}{2}a_2:1a_3$ and the indices (021). And by means of a side elevation the intercepts of the face (102) could also be determined.

19. COMPOSITE CRYSTALS AND CRYSTALLINE AGGREGATES

Loose, isolated crystals are comparatively rare in nature. They usually occur in groups. The grouping may be in parallel position (see Fig. 312), in the most irregular manner, or in the third condition of partial parallelism.

Twinning

The peculiar sort of grouping in partial hedrons in parallel position. parallelism is known as twinning; crystals so grouped are called twin-crystals. Many crystals are found to be composed of two parts, one-half of which apparently has been revolved 90° or 180° about a line called the twin-axis Other crystals have These may be called rotation twins. two portions symmetrically placed with reference to a plane MBTITHEOM: ALI VENNOBTICHEOM: AL

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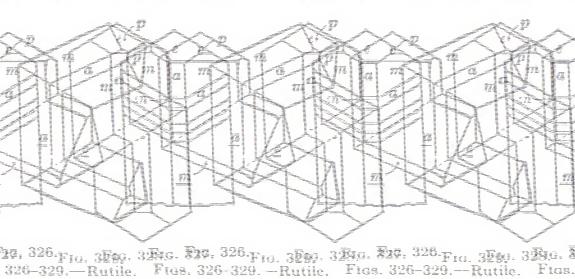
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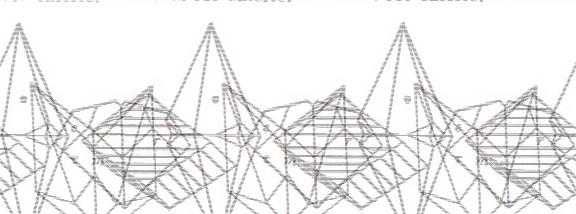
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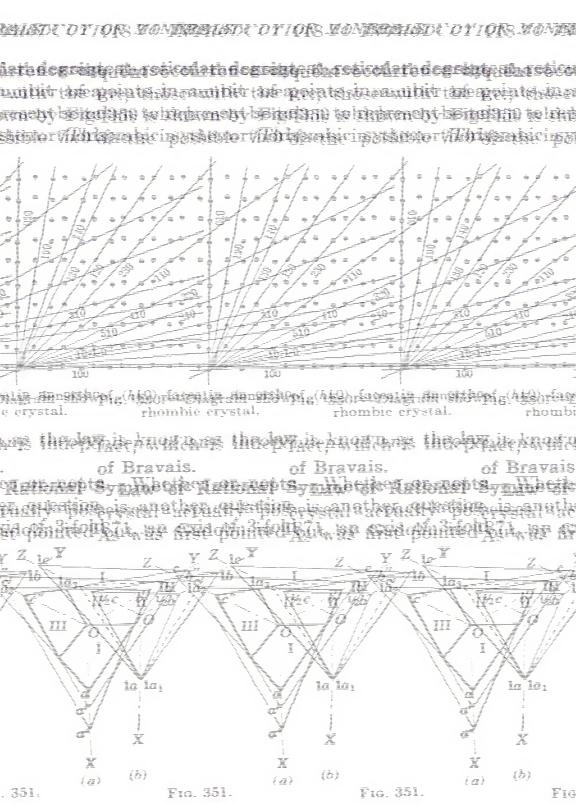
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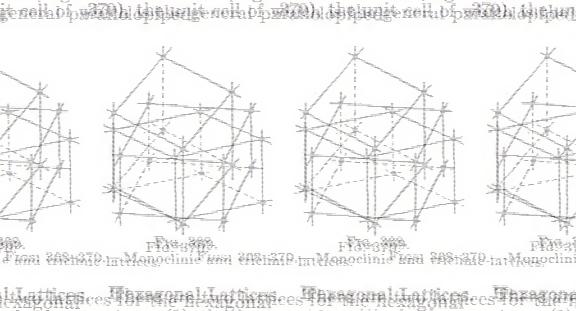
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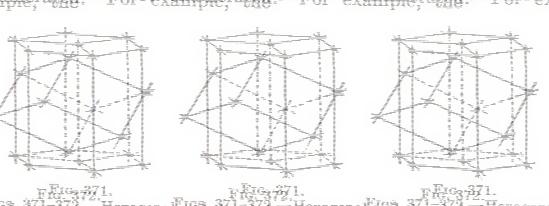
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The results obtained since that time have opened up one of the most interesting fields in the whole realm of science.

The first experimental work on crystal structure originated in an attempt to determine the nature of X-rays. It was doubtful whether X-rays consist of a wave-motion or are material in nature. Laue, a physicist at the University of Munich, conceived the idea that a crystal might act as a 3-dimensional diffraction grating for X-rays if the latter consisted of a wave-motion.

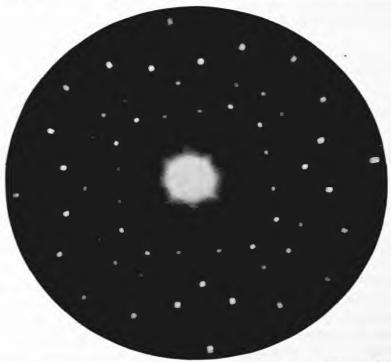


Fig. 375 .- Laue photograph of NaCl. (Photographed by Donald G. Hare.)

In order to test this, a beam of X-rays was directed upon a crystal plate of sphalerite (isometric ZnS), and a photograph was obtained which showed a central circular spot surrounded by elliptical spots of varying intensity arranged in a symmetrical manner similar to Fig. 375 which represents sodium chlorid.

The experiment had succeeded and proof was furnished that the mysterious X-rays discovered by Roentgen are the result of a wave-motion similar to, but with much shorter wave-length than, that of light. The central spot of the photograph is produced by the direct beam of X-rays and the other spots are due RTHERON (MINIERONGRIPERROUM) (MINIERONGRIPERROUM) (MINIERONG velipteriorpaterlangen degreelipteriorpaterlangen degreelipteriorpaterlangen degree uitteppetiolog (Thresholoidespoetiolog (Thresholoidespoetiolog (Ithresholoidespoetiolog)) verenner hedreald traditions en achterbaald traditions en achterbaald (milli

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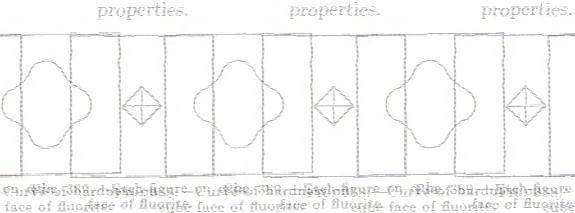
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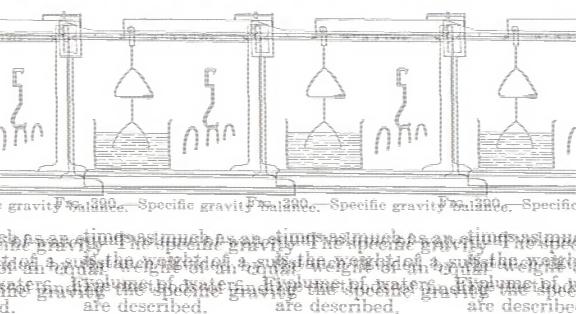
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Hammer. A small square-faced hammer of about 2 oz. weight is indispensable.

Anvil. A small block of well tempered steel, square or rectangular in cross-section, and about ½ in. thick, is convenient for powdering minerals and has many other uses.

Platinum Wire. No. 27 platinum wire is the best size for general use. The wire may be fused into a piece of glass tubing, or held in a special holder made for the purpose.



Fig. 477.—Portable blowpipe outfit. (Courtesy of the Denver Fire Clay Company.)

Test Tubes. The most convenient size is 12 cm. long and 15 mm. in diameter.

Glass Tubing. Soft glass tubing of 7 mm. outside diameter is best for most purposes, but it is well to have a variety of sizes. For some tests hard glass tubing is preferable.

Watch Glasses. These are needed especially for solubility tests. The best size is "half-deep" and 5 cm. in diameter.

Magnet. A magnetized knife blade answers the same purpose and is more convenient.

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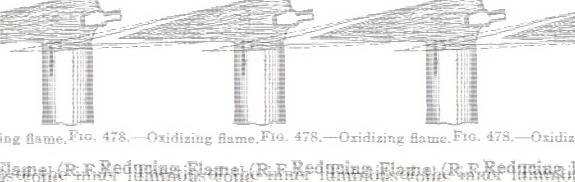
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observed, the only common well defined one is the octahedron. Figure 494 represents a typical crystal with grooved edges and triangular markings. Spinel twins of diamond are rather common.

The internal structure of diamond determined by X-ray analysis is like that of sphalerite (Fig. 383, p. 132) except that all the atoms are carbon atoms.

Cleavage perfect octahedral. This fact enables the diamond cutter to save considerable work.

H. = 10 (the hardest known substance). Color usually colorless or faintly colored, though brilliantly colored blue, green, and red stones are known. One variety known as carbonado is black and opaque.

Luster brilliant, known as adamantine. The rough uncut crystals have a peculiar

Fig. 494.—Diamond greasy appearance.

fraction is very high (n = 2.4175) for sodium light) which accounts for its brilliancy. The "fire" of the diamond is accounted for by its strong dispersion; the index of refraction for the red end of the spectrum is 2.402, while for the violet end it is 2.465. Diamonds are transparent to X-rays, while glass imitations are not.

Microscopic Tests. Fragments are irregular to triangular, dark between crossed nicols, and have high relief in methylene iodid.

Chemical Composition. Pure carbon. Upon heating the diamond in an atmosphere of oxygen it is converted into CO₂.

Blowpipe Tests. Infusible. Insoluble in acids.

Distinguishing Features. Diamond is distinguished from similar minerals by its superior hardness, its adamantine luster, and its comparatively high specific gravity.

The peculiarly rounded crystals with an apparently oiled

surface are unlike those of any other mineral.

Uses. On account of its great hardness, brilliancy, and rarity, diamond stands as the gem mineral par excellence. Among the famous historic diamonds are the Kohinoor, 186 carats, the

The international carat is 200 mg.

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SULFIDS 295

Name. Smaltite (smôl'tīt) was named by Beudant in 1832 since it was used in the preparation of smalt, a deep blue powdered cobalt glass used as a pigment.

MARCASITE GROUP—ORTHORHOMBIC

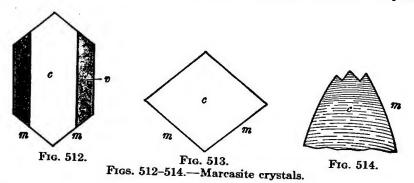
The following minerals: marcasite, FeS₂; arsenopyrite, FeAsS; gudmuntite, FeSbS; löllingite, FeAs₂; glaucodot (Co,Fe)AsS; safflorite, CoAs₂; and rammelsbergite, NiAs₂, constitute an isomorphous group parallel to the pyrite group. They are orthorhombic in crystallization, have a hardness of 5 to 6½, and are tin white to brass yellow in color. Only the first two of these minerals are considered, as the others are rare.

Marcasite

Orthorhombic (G. Markasit, F. Marcasite)

 FeS_2

Form. Marcasite occurs in orthorhombic crystals, in crystalline aggregates, and in rounded concretionary masses. Crystals



are usually tabular in habit and often elongated in the direction of the a-axis. Figures 512 to 514 represent typical crystals with the forms: $c\{001\}$, $m\{110\}$, $v\{013\}$. $mm(110:1\overline{10}) = 74^{\circ}$ 55'. Twins with $m\{110\}$ as twin-plane are common. A twin "fiveling" is represented in Fig. 325 (p. 110).

 $H. = 6 \text{ to } 6\frac{1}{2}.$ $G. = 4.9 \pm .$

Color pale brass yellow with a greenish tinge. Almost tin white when cleaned with dilute HCl (distinction from pyrite which is yellow). Luster metallic.

Chemical Composition. Iron disulfid, FeS_2 (Fe = 46.6 per cent). Analyses often show small amounts of arsenic.

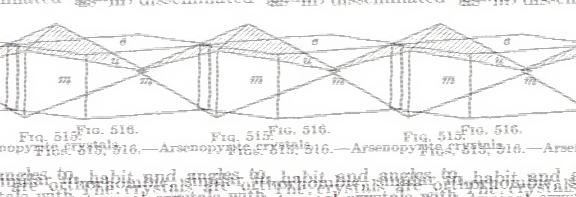
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mirror is formed. In the open tube on slow, gentle heating minute crystals of As₂O₃ are deposited and SO₂ also formed.

Soluble in HNO₃ with the separation of sulfur.

Distinguishing Features. Crystals of arsenopyrite resemble marcasite but are distinguished by the color of a fresh surface. Massive arsenopyrite resembles smaltite and often can only be distinguished by blowpipe tests.

Uses. Arsenopyrite is the chief source of the white arsenic (As₂O₃) of commerce. At Deloro, Ont., arsenopyrite is a gold ore.

Occurrence. 1. As an intermediate-temperature vein mineral associated with pyrite, chalcopyrite, galena, and other sulfids. Mother Lode of California.

- 2. In high-temperature veins.
- 3. In granite pegmatites.

Name. Arsenopyrite (är'sē-nō-pī'rīt) was named by Glocker in 1847. Before that time it had been known as arsenical pyrites and mispickel.

Sylvanite

Monoclinic (G. Schrifterz, F. Sylvanite) AuAgTe.

Form. Usually in branching crystal aggregates. Distinct euhedral crystals are rare. Twinned aggregates with {101} as twinning plane are characteristic.

Cleavage perfect in one direction parallel to b(010).

H. = 2. G. = 8.0 to 8.3.

Color silver white to steel gray with a yellow tinge. Luster metallic. Opaque.

Chemical Composition. Gold-silver tellurid, AuAgTe₄ (Au = 24.2, Ag = 13.3 per cent).

Blowpipe Tests. Easily fusible (at 1) on charcoal to a gray bead; at the same time it colors the flame bluish green and gives a white sublimate. With Na₂CO₃ the sylvanite is reduced to a pale yellow button.

Soluble in aqua regia with the separation of AgCl. Fragments of sylvanite dropped in hot conc. H₂SO₄ give a purplish-red coloration.

Uses. Sylvanite is an important ore mineral. It has been mined in Boulder Co., Colo. and near Kalgoorlie, Western Australia.

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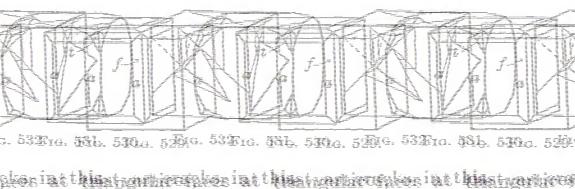
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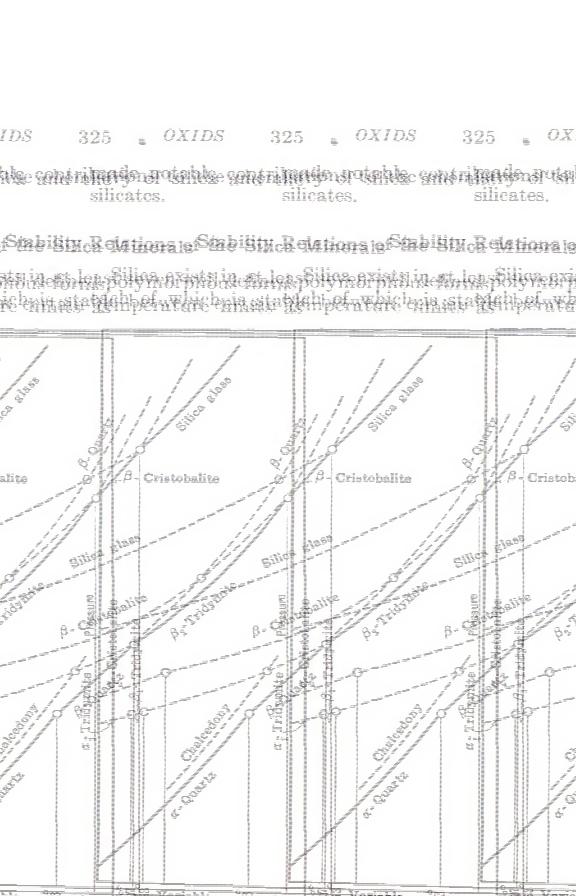
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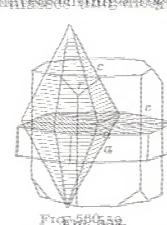
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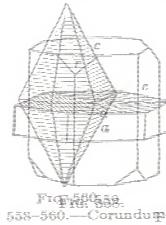
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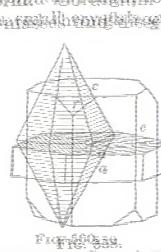
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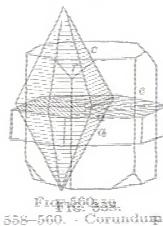
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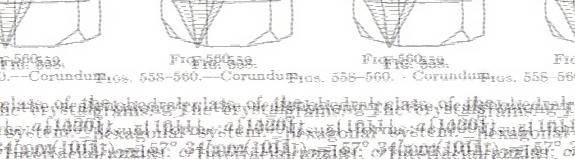
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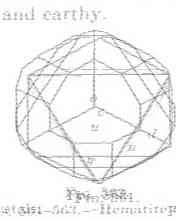
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Analyses of the Minerals of the Ilmen	tite Group
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	FeO	MgO	MnO	TiO ₂	Fe ₂ O ₃
Ilmenite Ilmenite Ilmenite (picroilmenite) Geikielite	22.4 36.5 24.4 6.3	0.5 0.6 14.2 28.5	0.3 2.7 	23.7 45.9 56.1 63.8	53.7 14.3 5.4 1.9

Blowpipe Tests. Infusible. The sodium carbonate fusion dissolved in HCl and boiled with metallic tin gives a violet solution.

Slowly soluble in HCl. Decomposed by fusion with KHSO₄.

Distinguishing Features. Ilmenite may be mistaken for hematite or magnetite. It fails to give the red streak of the former and the strong magnetism of the latter. When it is intimately mixed with hematite and magnetite, polished surfaces are necessary to identify it.

Occurrence. 1. As an accessory constituent of igneous rocks, especially diabases.

- 2. As a magmatic segregation in igneous rocks intergrown with magnetite and forming the so-called titaniferous magnetites.
- 3. As a prominent constituent of sands, especially the "black sands."

Name. Ilmenite (Il'měn-It) was named by Kuppfer in 1827 from the Ilmen Mountains, Russia.

RUTILE GROUP—TETRAGONAL

Cassiterite (SnO₂) and rutile (TiO₂) together with plattnerite (PbO₂) and polianite (MnO₂), are isomorphous; all are tetragonal dioxids of tetravalent metals.

CASSITERITE

Tetragonal (G. Zinnstein, F. Cassitérite)
(Ditetragonal dipyramidal class)

SnO₂

Form. Cassiterite is found in crystals, crystalline and colloform masses, pebbles, and grains (stream-tin). Crystals are

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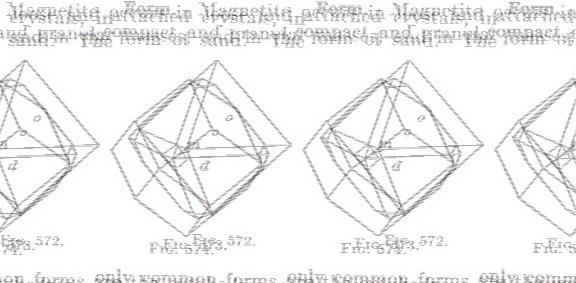
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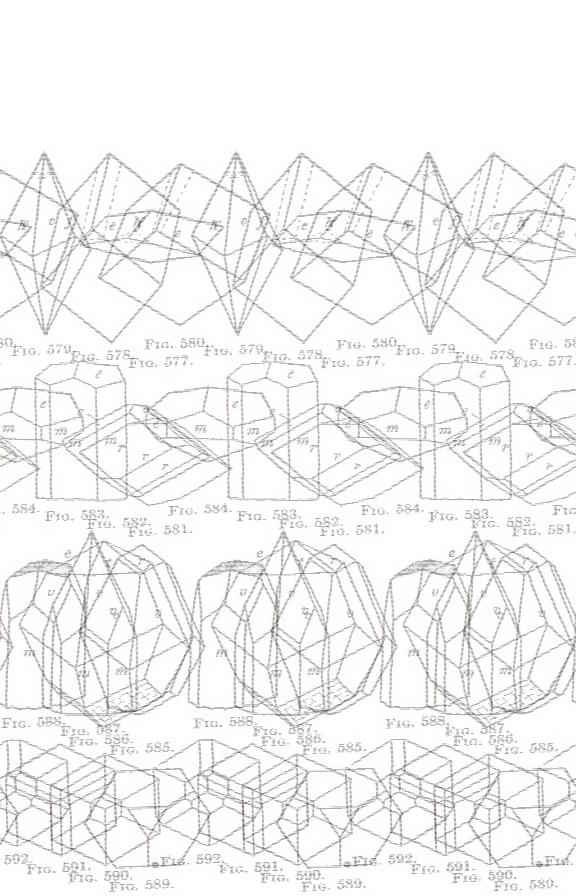
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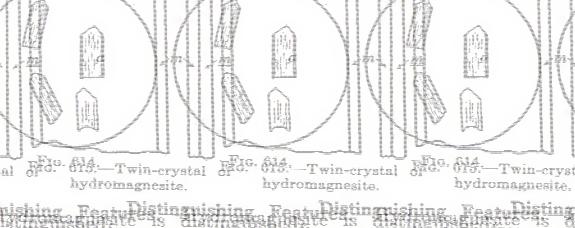
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prominent locality. Here the polyhalite zone is about 50 meters thick.

Polyhalite has also been found in abundance in the west Texas-New Mexico Permian basin. Here the associates are anhydrite, halite, sylvite, magnesite, kieserite (MgSO₄·H₂O), glauberite (Na₂SO₄·CaSO₄) and leonite (K₂SO₄·MgSO₄·4H₂O).

Name. Polyhalite (pŏl-ĭ-hā'līt) was named by Stromeyer in 1821 from the Greek *polus*, many, and *hals*, salt, in allusion to the number of salts in the mineral.

Brochantite

Orthorhombic (G. Brochantit, F. Brochantite) Cu4(OH)6SO4

Form. Brochantite is found in small prismatic crystals, in drusy crusts, and in fibrous masses.

Cleavage perfect parallel to $m\{110\}$.

H. = $3\frac{1}{2}$. **G.** = $3.9\pm$.

Color emerald green.

Optical Properties. Biaxial. Negative. $2V = 77^{\circ}$. $n_{\alpha} = 1.728$, $n_{\beta} = 1.771$, $n_{\gamma} = 1.800$. Birefringence strong. $n_{\gamma} - n_{\alpha} = 0.072$. Axial plane = (100). $c = \gamma$.

Microscopic Tests. Fragments are prismatic with parallel extinction.

Chemical Composition. Basic copper sulfate, Cu₄(OH)₆SO₄ or CuSO₄·3Cu(OH)₂ (Cu = 56.2, H₂O = 12 per cent).

Blowpipe Test. Fusible at $3\frac{1}{2}$. In the closed tube it turns black and yields water. It gives a green flame which is turned blue with HCl.

Insoluble in water. Soluble in HNO3 without effervescence.

Uses. At Chuquicamata, Chile, it is the principal ore of copper.

Distinguishing Features. Brochantite greatly resembles malachite and atacamite. Malachite effervesces with acid and atacamite gives a blue flame.

Occurrence. A secondary mineral associated with other copper minerals in the oxidized zone. Chuquicamata, Chile.

Name. Brochantite (brō-shănt'īt) was named by Lévy in 1824 in honor of Brochant de Villiers.

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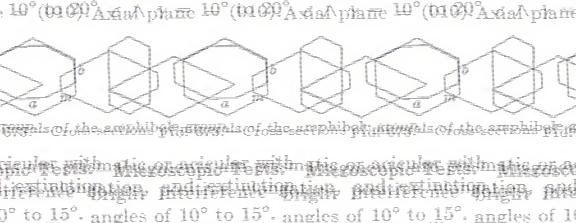
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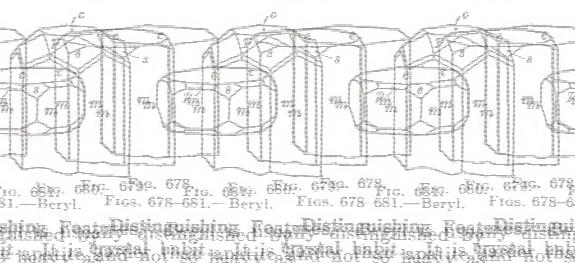
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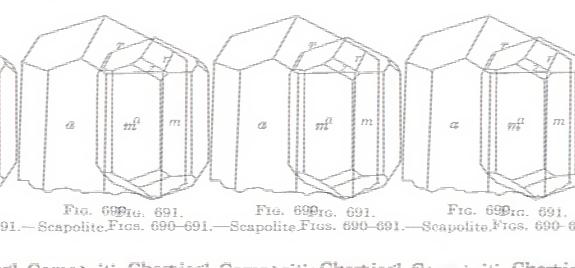
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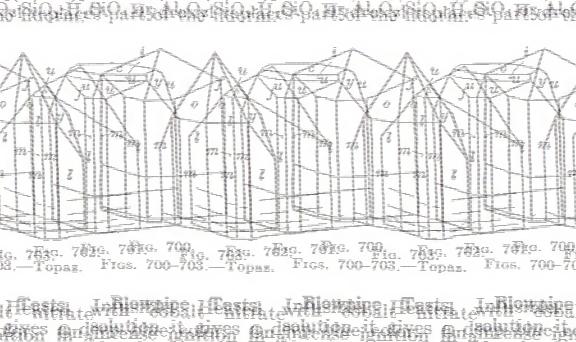
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Distinguishing Features. Chlorite is characterized by its perfect cleavage, absence of elasticity in the cleavage flakes, and green color.

Occurrence. 1. A secondary mineral in igneous rocks, formed by the alteration of such silicates as biotite, hornblende, augite, etc.

2. In schists, often forming independent rock masses, the chlorite schists. These are formed from original basic igneous rocks.

Name. Chlorite (klō'rīt) was named by Werner in 1789 from the Greek, chlōros, green, in allusion to the color of the minerals of the group.

Chloritoid

Monoclinic (G. Chloritspat, F. Chloritoïde) H2(Fe,Mg)Al2SiO7

Chloritoid is a member of a group of mica-like minerals called the *brittle micas*, which include clintonite, xanthophyllite, and margarite.

Form. Chloritoid usually occurs in disseminated plates or scales, which may be either single or aggregates. It is also found in foliated masses. Crystals are tabular pseudohexagonal in habit and are usually twinned.

Cleavage perfect in one direction parallel to $c\{001\}$, imperfect parallel to $m\{110\}$. Cleavage flakes are brittle.

H. = 6 to 7. G. = 3.4 to 3.6.

Color dark gray, greenish gray, dark green, to almost black. Subtranslucent. Luster pearly.

Optical Properties. Biaxial. Positive. $2V = 36^{\circ}$ to 60° . $n_{\alpha} = 1.715$ to 1.724, $n_{\beta} = 1.719$ to 1.726, $n_{\gamma} = 1.731$ to 1.737. Birefringence weak to moderate. $n_{\gamma} - n_{\alpha} = 0.007$ to 0.015. Axial plane 010. $c \wedge \gamma = 3^{\circ}$ to 21°. Pleochroic in greens or blues.

Microscopic Tests. Cleavage flakes are isotropic or almost so. Inclusions are usually present and some of these may have pleochroic haloes around them. The flakes give an eccentric biaxial interference figure.

Chemical Composition. Acid iron-magnesium-aluminum silicate, H₂(Fe,Mg)Al₂SiO₇ or (Fe,Mg)O·Al₂O₃·SiO₂·H₂O. (H₂O =

7.2 per cent for the pure iron end member.) Manganese may replace part of the iron. Manganian chloritoid is known as ottrelite. It is a prominent mineral in the schists in the Ardennes in Belgium.

Blowpipe Tests. Infusible in the closed tube at high tempera-

ture it yields water.

Insoluble in acids.

Distinguishing Features. Chloritoid is distinguished from the micas and chlorites by superior hardness and less perfect cleavage. The uncentred interference-figure of cleavage flakes is also a good distinction.

Occurrence. Chloritoid occurs as metacrysts in schists and phyllites. It is a product of low-grade metamorphism. Natick,

R. I., is a prominent locality for chloritoid.

Name. Chloritoid (klō'rĭ-toid) was named by G. Rose in 1837 on account of its resemblance to chlorite.

ANTIGORITE

Orthorhombic (G. Antigorit, F. Antigorite) H₄Mg₃Si₂O₉

Form. Antigorite has never been found in euhedral crystals, though it often occurs pseudomorphous after other crystallized minerals. It is usually compact or granular massive, but also occurs in fibrous, columnar, and lamellar forms. It is the main constituent of serpentine, which is properly considered a rock rather than a mineral.

H. = 3 to 4. **G.** = $2.5 \pm .$

Color green of various tints and shades, from greenish white to greenish black. It is also often yellow, brown, or red, and the color is not apt to be uniform, but is often in spots and streaks.

Optical Properties. Biaxial. Negative. $2V = 20^{\circ}$ to 90° . $n_{\alpha} = 1.560$, $n_{\beta} = 1.570$, $n_{\gamma} = 1.571$. Birefringence rather weak.

 $n_{\gamma} - n_{\alpha} = 0.011$. Axial plane = (100). $b = \alpha$.

Microscopic Tests. Fragments are irregular, or prismatic with parallel extinction and positive elongation. The interference colors are low first-order. The irregular fragments show aggregate structure between crossed nicols.

Chemical Composition. Acid magnesium silicate, H₄Mg₃Si₂O₉ or 2H₂O·3MgO·2SiO₂ (H₂O = 12.9 per cent). Part of the

magnesium is usually replaced by ferrous iron. Some analyses show a little aluminum and some a little calcium. These constituents are largely due to residual pyroxenes.

Analyses of Antigorite							
MgO	FeO	CaO	Al ₂ O ₂	Fe ₂ O ₃	SiO ₂	H ₂ O	Misc.
42.6	0.1	0.1		0.3	42.0	14.7	
36.5	1.9		5.1		42.9	13.2	NiO = 0.6
41.2	2.4				41.3	14.5	
36.8	7.2		2.6		41.6	12.7	$Cr_2O_1 = tr.$

Analyses of Antigorite

Blowpipe Tests. In the closed tube it gives water at a high temperature.

Decomposed by HCl with the separation of non-gelatinous silica.

Distinguishing Features. Antigorite is distinguished by its green color, moderate hardness (4), and its mottled, veined, or compact massive structure.

Uses. Serpentine is used as an ornamental stone.

Occurrence. 1. An alteration product of olivine and to a less extent of bronzite, forming the metamorphic rock serpentine from original peridotite.

- 2. A secondary mineral in seams and veins and on the border of serpentine rocks.
- 3. An alteration product of diopside and forsterite in crystalline limestones; these form the rocks known as ophicalcites.

Name. Antigorite (ăn-tīg'ō-rīt) was named by Schweizer in 1840 from the locality Antigorio Valley, Piedmont, Italy.

Chrysotile

Orthorhombic (G. Chrysotil, F. Chrysotile) H₄Mg₃Si₂O₉

Form. Chrysotile occurs in seams and always has a fibrous structure.

$$H. = 2\frac{1}{2}$$
. $G. = 2.2 \pm .$

Color green to golden yellow.

Optical Properties. Biaxial. Positive. $2V = 30^{\circ}$ to 35° . $n_{\alpha} = 1.542$, $n_{\beta} = ?$, $n_{\gamma} = 1.555$. Birefringence moderate. $n_{\gamma} - n_{\alpha} = 0.013$.

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indefinite in chemical composition to be considered minerals, while the hydrocarbons, though occasionally of definite composition, are organic. These two classes of substances may be treated in a supplement under the term other mineraloids.

VOLCANIC GLASS

Form. Amorphous, usually massive and structureless, but sometimes vesicular, spheroidal (perlitic), or banded.

$$H. = 5 \text{ to } 7.$$
 $G. = 2.2 \text{ to } 2.7.$

Color colorless, gray, black; sometimes green, brown, or red. Luster vitreous to resinous, dull if devitrified. Translucent to transparent on thin edges.

Optical Properties. Amorphous. n = 1.48 to 1.67 (increases in general with decrease of SiO_2).

Microscopic Tests. Fragments are irregular and are usually isotropic but may be doubly refracting due to strain, especially the perlitic varieties.

Chemical Composition. Variable, contains SiO₂, Al₂O₃, Fe₂O₃, FeO, MgO, CaO, Na₂O, and K₂O in amounts comparable to

Table	of	Analyses	of	Natural	Glass
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	SiO	Al,O,	FosOs	FeO	MgO	CaO	NazO	K ₂ O	H ₂ O	Misc.
Rhyolitic obsidian. Rhyolitic pitchstone. Andesitic perlite. Trachytic obsidian. Basaltic obsidian. Basaltic obsidian.	71.6 65.1 61.2	13.1 15.7 18.0 13.5	0.7 2.2 1.3 3.0	0.3 1.9 4.5 7.4	0.1 1.4 0.4 6.5	0.7 3.6 1.9 10.3	3.8 2.9 6.5 3.2	4.1 4.0 5.9 0.6	5.5 2.4 0.5 0.6	$MnO = 0.4$ $TiO_2 = 0.2$

those found in various types of igneous rocks ranging from the high silica content of rhyolite to the relatively low silica content of basalt. Water is low in obsidian, but comparatively high in pitchstone.

Blowpipe Tests. The fusibility of volcanic glass is characteristic. It is fusible to a vesicular enamel, but this enamel on further heating is infusible, which is due to the fact that the water is driven off. In the closed tube it gives more or less water (0.5 to 5 per cent).

Insoluble in acids. Tests for silica and the metals can be obtained by making a sodium carbonate fusion (see p. 239).

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Asphaltum or asphalt is a general name for a great variety of black, solid, more or less oxygenated hydrocarbons. They include besides the well-known Trinidad Lake asphalt, other varieties which have received special names such as albertite, gilsonite, grahamite, manjak, wurtzilite, and many others of local importance. Each of these has special characters of its own, but they are all similar, with a hardness of 1 to $2\frac{1}{2}$, specific gravity of 1.0 to 1.8. They melt easily and burn with a disagreeable odor. They are more or less soluble in alcohol, ether, turpentine, carbon bisulfid, and chloroform. The relative solubility in these various solvents is the best method of distinguishing the various kinds of asphaltums.

Asphaltum occurs in veins usually and rarely in lake deposits as on Trinidad Island. Asphaltum and semisolid hydrocarbons also occur as impregnation of sandstones or limestones. These bituminous sandstones and limestones have been used for paving in some parts of the United States.

The oil shales found so abundantly in western Colorado should also be mentioned here. On distillation they yield gas, crude oil, and ammonia. The crude oil on refining furnishes gasoline, burning oils, and paraffin wax.

Coal. Finally we have the coals which vary from lignite through sub-bituminous, bituminous, and semi-anthracite to anthracite. The chemical compounds in coal are for the most part unknown. Coal consists largely of carbon, hydrogen, and oxygen, with small amounts of nitrogen and sulfur. Analyses are usually given in a proximate form: moisture, volatile matter, fixed carbon, and ash.

The specific gravity of coal varies from about 1.2 to 1.7. The hardness reaches 2 to $2\frac{1}{2}$ in anthracite.

Coal is distinguished from other hydrocarbons by the fact that it is practically insoluble in the organic solvents (ether, alcohol, etc.).

Coal occurs in beds and so must be ranked as a rock as well as a mineraloid. Anthracite may be said to constitute a kind of metamorphic coal.

Coal is known to be of vegetable origin. All gradations have been traced from peat through lignite into various members of

the coal series. There is, however, disagreement as to the details of the formation of coal.

Salts of Organic Acids. The discussion of mineraloids would be incomplete without a brief reference to the salts of certain organic acids. One of these, calcium oxalate, occurs in plant tissues and is also found in coal beds in monoclinic crystals which have received the name whewellite. They have the composition $CaC_2O_4 \cdot H_2O$.

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(e) Gabbro-auganite Series.

The gabbro-auganite series includes igneous rocks with calcic plagioclase and some ferromagnesian mineral besides olivine as the essential minerals. The absence of olivine distinguishes this series from the olivine gabbro-basalt series.

Gabbro is here used for an olivine-free grained rock consisting essentially of calcic plagioclase (labradorite or bytownite) and augite, hypersthene, or hornblende.¹ The gabbros are usually coarser grained than diorites and the plagioclase is often altered to the mixture of minerals known as saussurite. The pyroxene may be altered to an amphibole called uralite.

Varieties rich in hypersthene are called **norite**. With increase in the plagioclase content the gabbros grade into **anorthesites** or labradorite rocks, which are sometimes used as an ornamental stone. With decrease in the plagioclase on the other hand, they pass into the pyroxenites, igneous rocks composed essentially of pyroxene. Quartz and orthoclase are present in some gabbros but such types are rare.

With decrease in the size of the mineral grains the gabbros grade through gabbro porphyries into auganite porphyries and auganites.

Auganite. This recently introduced term (A. N. Winchell, 1912) is a useful one for the dense equivalent of the olivine-free gabbro. Auganite differs from andesite in the fact that the plagioclase is more calcic than Ab₁An₁, and from basalt in the fact that olivine is lacking. It has usually been called augite andesite or olivine-free basalt, but the first name may be used for augite-bearing rocks with sodic plagioclase, and the name basalt may then be restricted to olivine-bearing rocks.

Leucite and nepheline occur in certain rare rocks of this group. One type, leucite tephrite, is abundant in certain parts of Italy.

(f) Olivine gabbro-basalt Series.

The olivine gabbro-basalt series includes igneous rocks with calcic plagioclase, olivine, and some other ferromagnesian mineral as the essential constituents. They contain less silica than the

¹ The distinction between diorite and gabbro is based upon the character of the plagioclase rather than upon the ferromagnesian mineral for it is known that in many subsilicic rocks hornblende has been formed from pyroxene at a late magmatic period.

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Nepheline and leucite occur in some rare types of basaltic rocks called basanite.

(g) Peridotite-limburgite Series.

The peridotite-limburgite series includes feldspar-free igneous rocks with olivine and pyroxene as essential constituents with their glassy equivalents. They are very low in silica and alumina and consequently high in lime, magnesia, and iron oxids.

Peridotites are grained rocks with olivine and pyroxene, but with little or no plagioclase. They grade into olivine gabbros on the one hand and into nearly pure olivine rocks (dunites) or nearly pure pyroxene rocks (pyroxenites) on the other hand. The pyroxene of peridotite may be augite, hypersthene, or enstatite, often an intergrowth of two of them. Hornblende is sometimes present. Common associated minerals are magnetite, ilmenite, chromite, and pyrrhotite. Garnet, spinel, and native iron are also found in peridotites.

Peridotites are seldom found in an unaltered condition but their alteration products, the hydrothermal metamorphic rocks known as serpentine, are very common. The principal mineral of serpentine is antigorite formed at the expense of the olivine and pyroxene of the original peridotite.

The peridotites are usually found in dikes rather than in subjacent rock masses. This is one reason why the petrographic classification of igneous rocks should be based upon texture rather than upon geological occurrence.

Limburgite, the dense equivalent of peridotite, is a very rare rock type and is mentioned here simply for the sake of completeness. It has been called magma-basalt. But the term basalt should be restricted to olivine-plagioclase-augite rocks. The limburgites contain phenocrysts of olivine and augite in a glassy groundmass. Augitite is used for the corresponding rock with augite alone.

OTHER FELDSPAR-FREE IGNEOUS ROCKS

Other feldspar-free rocks include some rare types in which the feldspathoids, nepheline and leucite, take the place of the feldspars. The leucitites, as the dense leucite-pyroxene rocks are called, are abundant in Italy. of MITHEORIGIN OF MITHEORIGIN OF MITHEORIGIN OF MITHEORIGIN CONTINUED CONTINUED AND ADDRESS OF THE OR A COMPANION OF MITHEORIGIN CONTINUED AND ADDRESS OF THE OR ADDRESS OF TH

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of two or more minerals such as graphic granite (quartz and feldspar) and perthite (microcline and albite), the variability in texture in different parts of the vein, and the presence of rare elements and rare minerals.

There are pegmatites corresponding to most of the known grained rocks but the granite pegmatites are especially common and the term pegmatite without any qualification usually refers to granite pegmatites. They are very prominent in New England, South Dakota, and southern California. The following minerals are found in granite pegmatites: quartz (both α -quartz and β -quartz, so that the temperature of formation was probably in the neighborhood of 573° C.), microcline (the common feldspar of pegmatites), orthoclase, albite, muscovite, lepidolite, tourmaline, amblygonite [(Li,Al)(F,OH)PO₄], topaz, beryl, apatite, fluorite, columbite, garnet, and also many rare minerals containing such elements as cæsium, zirconium, uranium, tantalum, etc.

In southern Norway over 70 minerals, most of them very rare, have been found in the syenite and nepheline syenite pegmatites of that region.

4. PYROCLASTIC ROCKS

The pyroclastic rocks are the fragmental products of igneous activity and their consolidated equivalents. They form a connecting link between volcanic igneous rocks, with which they are usually associated, and sedimentary rocks of mechanical origin. A volcanic ash worked over by water may form a tuffaceous sandstone.

The pyroclasts are formed during the explosive activity of volcanoes and thus these materials are often found interbedded with lavas. The fragments may vary in size from very large volcanic bombs through lapilli (from a few centimeters to a few millimeters in size) and so-called volcanic ash down to the extremely fine material known as volcanic dust, so fine that it floats in the atmosphere for hours or even days before it settles.

Consolidated volcanic ash is called tuff and consolidated rock of coarser fragments, agglomerate. Tuffs and agglomerates are common in volcanic regions. The older tuffs are firmly consolidated and on account of alteration often may be distinguished only with difficulty.

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quartz is undoubtedly the principal mineral of sand but rarely does the sand or gravel consist entirely of quartz. The common rock-forming minerals such as feldspar and mica are also found in sands and in favorable places where there has been natural concentration we have heavy minerals such as magnetite, ilmenite, garnet, zircon, rutile, tourmaline, and occasionally diamond, gold, platinum, and cassiterite. Locally almost any mineral may occur. For example, on the Hawaiian beaches olivine is the only common mineral of the sand. In the arid regions of certain parts of New Mexico and Utah the dune sands are made up entirely of grains of gypsum. On the island of Ceylon there are found gem-bearing gravels with such minerals as spinel, corundum, garnet, zircon, ilmenite, monazite (Ce,La,-Di)PO₄, and chrysoberyl (BeAl₂O₄).

Consolidated

Breccia. A rock made up of coarse angular fragments is known as a breccia. The fragments may be quartz, chalcedony, calcite, or almost any mineral or even rock fragments. The cementing material may be the same as, or different from, the fragments.

Conglomerate. Breccias grade into conglomerates, the fragments of which are subangular to rounded and are due to mechanical wear. The pebbles of conglomerates usually have been transported considerable distances, while the angular fragments of breccias have been cemented before much mechanical action took place.

Sandstones. Consolidated sands are called sandstones. The grains are subangular to rounded. The principal constituent of sandstone is quartz, but other minerals such as feldspar, muscovite, and calcite are not rare. A sandstone with appreciable amounts of feldspar is known as arkose. Closely related to arkose but containing in addition to quartz and feldspar miscellaneous minerals and rock fragments such as bits of shale is the rock known as graywacke. It is a kind of fine-grained conglomerate. According to the character of the cementing material sandstones are classed as calcareous, ferruginous, siliceous, or bituminous. On account of porosity fossils are comparatively rare in sandstones. The calcite in calcareous sandstones is due in part to a recrystallization of the calcium carbonate of calcareous

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Limestone. The most common rock of organic origin is limestone, which consists almost entirely of the mineral calcite, usually in a massive form. While often apparently amorphous it is really microcrystalline. Calcium carbonate is present in most natural waters and is secreted by various organisms such as molluscs, brachiopods, echinoderms, bryozoa, corals, certain sponges, and some foraminifera. The organisms just named are animals but some plants such as algæ also secrete calcium carbonate. Most of the limestones are marine in origin but fresh water limestones are also known. They may be recognized by the presence of plant stems and fresh water mollusc shells such as those of snails.

When the organisms die, the organic material decays and there remain the shells and tests and other hard parts such as crinoid stems and coral fragments. They accumulate and are more or less broken up, often ground to bits and thus form a calcareous mud which on hardening produces a limestone.

Besides the mineral calcite, sedimentary limestones may contain dolomite, quartz, chalcedony, and such minerals as barite, celestite, anhydrite, gypsum, siderite, aragonite, collophane, and pyrite. These minerals occur for the most part in seams and cavities along with recrystallized calcite.

Varieties of limestone are based in part upon the character of the fossils present, e.g., crinoidal limestone, coral limestone, nummulitic limestone, and shell limestone (the shell limestone on the coast of Florida is called coquina and represents the first stage in the formation of some limestone). Chalk is a very fine-grained porous limestone formed by the consolidation of calcareous ooze which is made up principally of tests of foraminifera. Though chalk is apparently amorphous, microscopic examination proves it to be microcrystalline.

Depending upon the impurities present such varieties of limestone as siliceous, arenaceous, argillaceous, and bituminous are recognized. These show gradations toward cherts, sandstones, shales, and hydrocarbons respectively. A limestone that will take a polish, whether sedimentary or metamorphic, is called marble. of MIXHKAQRIGIN of MIXHKAQRIGIN of MIXHKAQRIGIN of MIXHKAQRIGIN

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the calcium and the substitution of magnesium in its place, is called dolomitization. But in some cases dolomitic limestones have been formed by the replacement of limestone after it was uplifted above the sea, sometimes by the agency of meteoric waters and sometimes by hydrothermal solution in connection with ore deposition.

Chert. Most of the compact massive rocks made up largely of chalcedony and known as chert or flint were probably formed by the replacement of limestones or shales by means of siliceous solutions. Although radiolaria, diatoms, and sponge spicules doubtless have contributed their share of silica, most of the fossils present in cherts were originally calcareous. Some of the cherts, however, may have been formed from radiolarian ooze.

Phosphorite or so-called phosphate rock is a sedimentary rock composed largely of the amorphous mineral collophane, a calcium carbonate-phosphate. The associated minerals include dahllite (the crystalline equivalent of collophane), calcite, fluorite, gypsum, chalcedony, and quartz. Although the phosphatic material is in part of organic origin, most of the phosphorites seem to be replacement of limestones as fossils present were originally calcareous.

The high-grade phosphorites are used extensively as a source of calcium superphosphate, which is employed as a fertilizer. They occur in Florida, Tennessee, and also in southeastern Idaho and the adjoining portions of Wyoming and Utah.

6. METAMORPHIC ROCKS

The third great group of rocks consists of those which though originally either igneous or sedimentary now possess characters that entitle them to recognition as a separate group. New minerals or new textures or both have been developed by geologic agents such as heat, pressure, or hydrothermal solutions. The term metamorphism, though sometimes used in a broad sense for practically all rocks except the igneous, is generally restricted to changes brought about by hypogene or deep-seated forces acting from within the earth's crust or outer shell. Used in this sense metamorphism excludes weathering and the replacement of sedimentary rocks by solutions of meteoric water.

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Hypogene and supergene chalcocite may usually be distinguished by a careful microscopic examination of polished surfaces.

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ZIRCONIUM

Zircon, ZrSiO4

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Ā

A. Symbol used for an axis of n-fold symmetry.

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Accessory minerals of igneous rocks, 527

Achroite. A colorless variety of elbaite.

Acicular. Needle-shaped.

"Acid" igneous rocks. The same as persilicic rocks, 526

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Actinolite. Ferriferous tremolite.

Acute bisectrix, 181

rhombohedron, 87

Adamantine luster, 143

Addition and subtraction rule, 32

Adjustments of the polarizing microscope, 162

Adularia, 430

Aegirine, 448

Agate. A banded or variegated variety of chalcedony.

Agglomerate, 540

Aggregate polarization. The intricate structure of microcrystalline substances between crossed nicols.

Alabandite, 285

Alabaster. Translucent massive gypsum. Formerly used for onyx marble.

Alaskite, 532

Albite, 434

twinning, 433

Alexandrite, 345

Alkali feldspar. A collective name for orthoclase, microcline, and albite.

Allanite, 478

Allotriomorphic. The same as anhedral.

Almandite, 461

Alpha (α). (1) The angle between the b- and c-axes of reference in the triclinic system. (2) The direction of the fastest ray in anisotropic crystals.

Alteration of minerals, 524

Alumina. Aluminum oxid (Al₂O₄).

Aluminates, 339

Aluminum minerals, 593

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Alundum. Artificial Al₂O₃ used as an abrasive.

Alunite, 400

Amalgam, 263

Amazon-stone. The green variety of microcline.

Amber, 517

Amblygonite, 418

Amethyst. The purple variety of quartz often used as a gem.

Amorphous condition, 5

Amphibole. A group of silicates including anthophyllite, tremolite (actinolite), bornblende, baseltic hornblende, and glaucophane, 450

Amphibolite, a non-schistose metamorphic rock made of hornblende

Amygdaloidal, 528

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Analcite. Synonym of analcime.

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DIOPSIDE, 444

Dioptase, 476

Diorite, 535

Diploid, 104

Diploidal class, 103

Dipyramid, 20

Directions image, 181

Dispersion. The divergence of the optical constants for different parts of the spectrum.

Disphenoid, rhombic, 22

tetragonal, 22

Disseminated. Scattered through a rock or vein in small quantities.

Disthene. Synonym of kyanite.

Ditetragonal dipyramidal class, 67 prism, 20

pyramidal class, 74

Ditrigonal dipyramidal class, 81 prism, 20

pyramidal class, 93

Dodecahedron, 97

Dodecants. The twelve divisions into which space is divided by the four axes of reference of the hexagonal system.

Dog-tooth spar. A variety of sharp-pointed calcite crystals.

DOLOMITE, 360

Dolomitic limestone, 549

Dolomitization, 550

Domatic class. The crystal class with a single plane of symmetry, 62

Dome, 19

Double refraction, 158

Dravite, 485

Drawing of crystals, 41

Drusy. Apparently sprinkled over with minute crystals.

Dry-bone. A local synonym of smithsonite (Wis.) or cerussite (Mo.)

Dumortierite, 487

Dunite, 538

 \mathbf{E}

Eclogite. A metamorphic rock containing garnet and omphacite, a variety of pyroxene.

Edge of a crystal, 6

Effervescence. The bubbling caused by the evolution of gas such as CO₂.

Efflorescent. Gives up its water of crystallization on standing.

Eglestonite, 287

Elæolite. A synonym of nepheline. Elastic. An elastic mineral springs back when bent in contrast to a flexible mineral which remains bent.

Elbaite, 485

Electrum. A naturally occurring alloy of gold and silver with over 20 per cent. of silver.

Elements, axial. A collective name used for the axial ratio and the angles between the axes of reference.

list of chemical, 198 occurrence of chemical, 197 of symmetry, 15

Ellipse, optic, 177

Ellipsoid, optic, 179

Emerald. The clear green variety of beryl used as a gem.

Emery. A mixture of corundum with magnetite.

Enantiomorphous, 24

Enargite, 307

Endlichite, 415

Enstatite, 443

EPIDOTE, 479

Epsomite, 395

Equant. The same as equidimensional.

Equation of zone control, 32

Erubescite. A synonym of bornite.

Erythrite, 294

Essonite (Hessonite). A variety of grossularite garnet used as a gem.

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916). calcite. (1851–1

Indicatrix. A synonym of the index ellipsoid.

Indices, Miller, 27

of refraction, 148

Indicolite. A deep blue variety of elbaite.

Indigo copper. A synonym of covellite.

Injected igneous rocks, 529

Integrity of zone character, 35

Interaxial angles, 39, 40

Intercepts, 25

Interfacial angle, 7

Interference, 163

colors, 163

of light, 163

Interference figures, 181 color chart, 168, 170

Intergrowth. An interlocking arrangement of two substances produced by simultaneous crystallization.

Internal structure of crystals, 116

Intrusive, 529

Intumesce, 225

Inversion, 14

twin, 109

Iodid flux, 214

Iodyrite, 84

Iolite. Synonym of cordierite.

Iridescence. A rainbow effect produced by the interference of light in thin surface films.

Iridosmine, 263

Iron, 273

minerals, 596

tests for, 235

pyrites. A synonym of pyrite. Isochromatic curves. The colored curves of interference figures in white light which are curves of equal retardation.

Isodimorphism, 355

Isogyres. Those portions of interference figures that are dark when viewed in white light.

Isometric system, 95

Isomorphism, 205

Isomorphous mixtures, 207 Isotropic, 177

J

Jacobsite, 339

Jade. A tough compact green or greenish-white ornamental stone consisting either of an amphibole (nephrite) or of a pyroxene (jadeite).

Jadeite, 449

Jamesonite, 301

Jargon. A pale-colored variety of zircon used as a gem.

Jarosite, 401

Jasper. A red to yellow or brown variety of chalcedony or quartz colored by iron oxids.

Jasp-opal. Opal colored yellow, red, or brown by iron oxids.

Jeffersonite, 448

Jolly balance. A specific gravity balance in which use is made of a spiral brass wire.

K

Kainite, 390

Kaolin. Impure kaolinite.

Kaolinite, 502

Kaolinization. The process by which kaolinite is formed.

Kernite, 377

Kidney ore. A colloform variety of hematite from England.

Kieserite, 203, 392

Kimberlite, 265

Kunzite. A transparent lilaccolored variety of spodumene used as a gem.

Kyanite, 475

L

λ (lambda). Symbol used for wavelength of light.

Labradorite, 437 Laccolith, 529 Lamellar, 112 Lampadite. A cupriferous variety of psilomelane. Lanthanum minerals, 597 Lapilli, 540 Lapis lazuli. A mixture of haüyne. calcite, diopside, etc. Latite, 534 Laue, M., German mathematical physicist. Laue photographs, 127 Laumontite, 513 Lautarite, 548 Lava, 529 Law of constancy of interfacial angles, 16 of maximum reticular density, 117 of rational indices, 25 of rational symmetric intercepts, 118 Lawsonite, 483 Lazulite, 420 Lazurite, 442 Lead minerals, 597 tests for, 236 Lechatelierite, 324 Left-handed quartz crystal, 318 Length-fast, 176 -slow, 176 Lenticular. Lens-shaped. Leonite, 392 Lepidolite, 493 Leucite, 439 Leucitite, 538 Light, convergent, 181 nature of, 145 polarized, 156 Limburgite, 538 Lime. Calcium oxid (CaO), often incorrectly used for calcium carbonate. Limestone, metamorphic or crystalline, 552 oölitic, 548 sedimentary, 544

Limit form, 36 LIMONITE, 349 Linarite, 393 Lindgren, W., American geologist (1860----), Linear projection, 30, 47 Lithia. Lithium oxid (Li₂O), Lithia mica - lepidolite, 493 Lithium minerals, 597 tests for, 236 Lodestone. A variety of magnetite which acts as a magnet. Löllingite, 295 Lussatite. An intimate mixture of fibrous chalcedony with opal. Luster, 142

M Macle. Synonym of a twin-crystal. Macroscopic. The same megascopic. Magma, 526 Magnesia. Magnesium oxid (MgO) Magnesian limestone. Limestones containing magnesium, but in quantities insufficient for dolomite. Magnesioferrite, 339 Magnesite, 363 Magnesium minerals, 597 tests for, 236 Magnetic iron ore. A synonym of magnetite. pyrites. A synonym of pyrrhotite. MAGNETITE, 341 MALACHITE, 372 Malformed crystals, 16 Malleable. Capable of being hammered out flat. Maltha, 517 Mammillary, 113 Manebach twin, 429 Manganese epidote. Synonym of piedmontite. Manganese minerals, 598

tests for, 237

Manganite, 349

Manganophyllite, 482

Marble. Any limestone that will take a good polish, but also used by some for a metamorphic limestone.

Marcasite, 295

Margarite, 497

Marmatite. Ferriferous sphalerite.

Martite. A pseudomorph of hematite after magnetite.

Massive. Without definite form or structure.

Measurement, of axial angle, 187 of crystals, 7 of index of refraction, 151 of interfacial angles, 7

Mechanically formed sedimentary rocks, 541

Mediosilicic igneous rocks, 526

Meerschaum. Synonym of sepiolite.

Megascopic. Capable of being seen with the unaided eye in contrast with microscopic.

Meigen's test, 368

Melaconite, 558

Melanite. A black titaniferous variety of andradite.

Melanterite, 396

Melatopes. The dark spots or "eyes" of biaxial interference figures. They represent the emergence of the optic axes.

Menaccanite. A synonym of ilmenite.

Mercury (mineral), 268

Mercury minerals, 598

tests for, 237 Merohedral, 40

Merosymmetry, 40

Merwin's flame-color screen, 213

Mesitite, 363

Metacolloids. Microcrystalline substances of colloidal origin.

Metamorphic rocks, 550 limestone, 552

Metamorphism, 550

Metasilicates, 427

Metasomatic replacement. The replacement of a rock mass by solutions.

Methylene iodid, 139

Miarolitic. A term applied to the structure of injected rocks containing cavities lined with euhedral crystals.

Mica group, 491

plate, 175

Micaceous, 112

Microchemical analysis, 243

gypsum, 250 struvite, 253

MICROCLINE, 431

Microcosmic salt. The same as salt of phosphorus or acid sodium ammonium phosphate used in blowpipe analysis.

Microlites. Minute crystals found in volcanic glasses.

Microperthite. Perthite on a small scale.

Microscope, polarizing, 160

Miller, W. H., English crystallographer and mineralogist (1801 -1880).

Miller indices, 27

Millerite, 288

Mimetic twinning. The tendency of twinning to raise apparently the grade of symmetry and thus imitate other crystals.

Mimetite, 413

Mineral, definition of, 1

Mineral species, 261

Mineraloid, definition of, 262

Minette, 350

Mirabilite, 548

Mispickel. A synonym of arsenopyrite.

Moissanite, 265

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Molybdates, 402

MOLYBDENITE, 279

Molybdenum minerals, 598 tests for, 237

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Onyx. A variety of chalcedony with sharply contrasted bands of colors.

Onyx, Mexican. The same as onyx marble.

Onyx marble. A banded variety of calcite or aragonite formed from water solutions.

Oolite. A rock made up of minute spheres formed by concretionary action.

Ooliths. The individual spheres of an oolitic rock.

Oolitic limestone, 548

structure. The structure of an oolite.

OPAL, 321

Opalescence. The peculiar milky appearance often seen in opal.

Opalized wood. A replacement of wood by opal.

Open form, 24

Open tube tests, 219

Ophicalcite, 555

Optic axes, 180

axis. The c-axis in the tetragonal and hexagonal systems.

ellipse, 177

ellipsoid, 177

Optic normal, 178

Optical anomalies, 193

character, 178

constants, 180

indicatrix. The same as the index ellipsoid.

orientation, 193

properties, 145

sign, 178

tests, 191

Optics, crystal. Optical crystallography or the division of crystallography that deals with the transmission of light in crystals.

Order of succession, 522

Ordinary ray, 158

Ores, 556

Organically-derived sedimentary rocks, 543

Oriental amethyst. A purple variety of corundum.

emerald. A green variety of corundum.

ruby. The true ruby, a variety of corundum.

sapphire. The true sapphire, a variety of corundum.

topaz. A yellow variety of corundum.

Orientation of a crystal. The placing of a crystal in its conventional position with the c-axis vertical and the a-axis pointing toward the observer.

optical, 193

Origin of minerals, 521

Orpiment, 276

Ortho-axis. The b-axis in the monoclinic system.

ORTHOCLASE, 428

Orthogonal zone, 33

Orthographic drawing or projection,

Orthorhombic system, 53

Orthosilicates, 427

Ottrelite, 498

Oxidizing flame, 217

Oxids, 317

Oxyhornblende. Synonym of basaltic hornblende.

Ozocerite, 517

P

P. Symbol used for plane of symmetry.

Pallasite, 464

Paragenesis, 522

Parallel growth, 108

extinction, 171

Parameters. The relative intercepts of the unit face (111) on the axes of reference.

Paramorph, 524

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Salt. (1) Synonym of halite. (2) Compounds formed by the union of bases with acids.

Salt of phosphorus. Hydrous acid sodium ammonium phosphate used in bead tests.

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Sand, 541 Sandstone, 542

Sanidine. A clear transparent variety of orthoclase found in volcanic igneous rocks.

Sapphire. The transparent blue variety of corundum.

Sard. A brownish-red variety of chalcedony used as a gem.

Sardonyx. Agate with red and white bands.

Sassolite, 539

Satin-spar. A fibrous variety of gypsum.

Scalar, 136

Scale of fusibility, 225 hardness, 141

Scalenohedral, hexagonal, class, 86 tetragonal, class, 71

Scalenohedron. There are two kinds of scalenohedrons but when no qualifying term is used the hexagonal scalenohedron is meant.

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Schistose. With the characters of a schist.

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Schoenflies, A. M., German mathematician (1853-1928).

Schorl. An old name for tourmaline.

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enrichment, 559

Sectile. Capable of being cut by a knife but not flattened out by a hammer.

Sedimentary rocks, 541

Selenite. The cleavable variety of gypsum.

Selenite plate. A plate of selenite showing the sensitive violet (q.v.).

Semi-opal. Common opal as distinguished from precious opal and fire opal.

Sensitive violet. The purple interference color between red of the first-order and blue of the second-order, 172

Sepiolite, 505

Sericite. A finely crystalline variety of muscovite.

Serpentine (rock), 554

minerals of. Antigorite, chrysotile.

Serpentinization. The alteration of peridotite or related rock to serpentine.

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glance. A synonym of argentite.

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Yü. The Chinese name for jade.

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